The Distribution of Glaciers in the American Rocky Mountains

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SHORT NOTES

THE DISTRIBUTION OF GLACIERS IN THE ROCKY MOUNTAINS OF THE UNITED STATES

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ABSTRACT. Evidence from aerial photographs, maps, and field checks indicates that 319 glaciers lie in cirques of the Rocky Mountains, south of the United States–Canadian border. On a subcontinental scale, the distribution of glaciers is highly clustered, with larger and denser clusters located in the northern Rocky Mountains. Lesser concentrations of small glaciers occur in the southern Rocky Mountains. The total area of glaciers in the Rocky Mountains of the U.S.A. is 78.9 km².

RESUME. La distribution des glaciers dans les Rocky Mountains des États-Unis. Photographies aériennes, cartes, observations de terrain montrent que 319 glaciers occupent des cirques dans les Rocky Mountains au Sud de la frontière États-Unis–Canada. À une échelle subcontinentale la distribution des glaciers est largement groupée avec des groupes plus grands et plus denses dans le Nord des Rocky Mountains. De moindres concentrations de petits glaciers se produisent dans le Sud des Rocky Mountains. Le total de la zone englaciée dans les Rocky Mountains des États-Unis est de 78,9 km².


INTRODUCTION

Glaciers are an integral part of many alpine landscapes in the Rocky Mountains. Most of the alpine landscapes have been shaped by predecessors of the present ice masses, while present locations and distributions of the glaciers exert profound influences on the hydrologic systems of the mountains and the water supply. An understanding and appreciation of the geography of the Rocky Mountains would be incomplete without an analysis of the distribution of glaciers, but lack of an adequate data base has hampered a complete census of ice bodies. This note presents evidence from recent aerial photography, maps, and field checks towards an account of glacier distributions south of the United States–Canadian border.

GLACIER CENSUS

Research workers have been documenting glaciers and their activity in the Rocky Mountains for almost a century. Three attempts have been made at a complete census of all glaciers in the Rocky Mountains. The earliest scientific reports were by Hayden (1878[a], 1878[b]) and Holmes (1878). Russell’s early work focussed on a few well-known areas of the Rocky Mountains, but made no attempt at a complete census (Russell, 1885, 1897). The monumental work of Dyson (1952) was the first effort to collate aerial photograph references for all glaciers of the Rocky Mountains, but many of the photographs were small scale and of poor quality. Dyson frequently relied on second-hand reports from untrained observers, such as forest rangers, for supporting evidence. Meier (1961) later refined the glacier count, but some field experience of the present author has revealed inadequacies in his accounting. Denton’s general work is the most complete account to date, but provides little new information and perpetuates previous errors (Denton, 1975). Since the time of the major studies, new, high-quality, large-scale photography has been made available by the U.S. Forest Service and the U.S. Geological Survey. This is supplemented by numerous topographic maps with scales as large as 1:24 000.

In the present study, glaciers were initially identified and outlined for study on U.S. Geological Survey or U.S. Forest Service aerial photographs at scales ranging from 1:15 000 to 1:31 000.

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Selection of the photographs is crucial, particularly with respect to date, because snow cover lingers late in spring and early summer. The last week in August or the first week in September offers the best chance for good aerial photos for the alpine areas of the Rocky Mountains because early fall storms leave snow in mid-September. Even during this later period, it is often difficult to differentiate between snow banks and glaciers, but field checks have revealed three primary lines of evidence: visible ice, evidence of movement, and crevasses.

In many cases, glacial ice is visible in the photographs; this resolves all questions of identification. Glacial ice is darker than surrounding snow, and at scales from 1:15,000 to 1:31,000 ice usually has a coarser texture than powdered or granular snow. Evidence of past movement usually consists of concentric morainic ridges about the lower end of the glacier. Protalus ramparts occasionally present a moraine-like appearance, but they occur much closer to the cirque or valley wall than do true moraines. Ramparts are usually sharp-crested crescents that may be lined end to end, but they rarely form concentric rings (Daly, 1912; Richmond, 1962). Crevasse patterns provide the final clues in the differentiation between small glaciers and snow banks. In the small glaciers of the Rocky Mountains, crevasses are almost always sharply defined transverse cracks with noticeable curvature. Such crevasses may occur anywhere on the glaciers, but most commonly they form bergschrunds near the tops of glaciers, and calving crevasses near the bottoms. Snow banks may develop bergschrunds in rare cases, but compaction prevents crevasses from forming near the bottom. Rapid glacier retreat during the early 1900's has made the following situation common: a steep glacier front (frequently with crevasses), with a small evacuated basin immediately down valley, and a recent moraine built before 1850 that dams a small lake of melt water. Permanent snow banks do not normally occur in a similar situation unless the pre-existing glacier has completely melted and has been replaced by permanent snow, an unlikely circumstance given the relatively small magnitude of climatic changes in the last few decades.

Table I and Figure 1 show the census of glaciers in the Rocky Mountains as determined from aerial photography and ground checks. For the purpose of this study a glacier is defined as a body of ice, with a surface area equal to or greater than 0.05 km² (about 11.5 acres), originating on land from recrystallization of solid precipitation, and showing evidence of past or present movement. Substantive parts of the definition are borrowed from Flint (1971), but the size limit is designed to include ice bodies traditionally considered in the literature as "glaciers".

The mean shapes of the glaciers were calculated by measuring the length from highest to lowest points and then dividing this value by the maximum width. Measurements were made photogrammetrically with ground checks in questionable areas. A shape value greater than 1.0 indicates that the glacier is longer than it is wide, forming a tongue of ice. A shape value less than 1.0 indicates a glacier that is wider than it is long, while a value close to 1.0 is typical of many cirque glaciers.
The three major concentrations of glaciers in the Rocky Mountains south of the United States–Canadian border (the Livingstone and Lewis Ranges in Glacier National Park, the Beartooth Range in southern Montana, and the Wind River Range in west-central Wyoming) account for more than 75% of the total number of glaciers and 89% of the total area covered by glacial ice. The total number of glaciers given here exceeds earlier estimates by Dyson and Meier partly because of the photographic coverage which is more complete and is of better quality than that available to previous workers. The total glacierized area exceeds the Meier (1961) estimate by about 15%.

The largest glaciers are concentrated in the Wind River Range, which contains 20% of the glaciers by number, but 40% by area. This result may be contrasted with that for the Front Range, which has 7% of the glaciers, but less than 2% of the glacierized area. Glaciers in the Front Range are small. This is apparently a response to the climatic conditions associated with the low latitude of the range. Although glaciers generally become more numerous as distance from the equator increases, the distribution of glaciers in the Rocky Mountains province is highly clustered with the larger and denser clusters occurring in the north.

Part of the explanation of the distribution of glaciers in the Rocky Mountains is concerned with air-mass domination. All existing glaciers in the region are inside the area occupied by mild Pacific air masses for more than 50% of the time (as mapped by Bryson and Wendland, 1967). The southern boundary of the glacier zone coincides with the boundary between the Pacific-dominated area and regions more strongly affected by the warm dry continental air originating in the interior south-western United States.

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